

Meso- γ scale Water Vapor Distributions of the Rainfall Systems Estimated by the Tomography Method

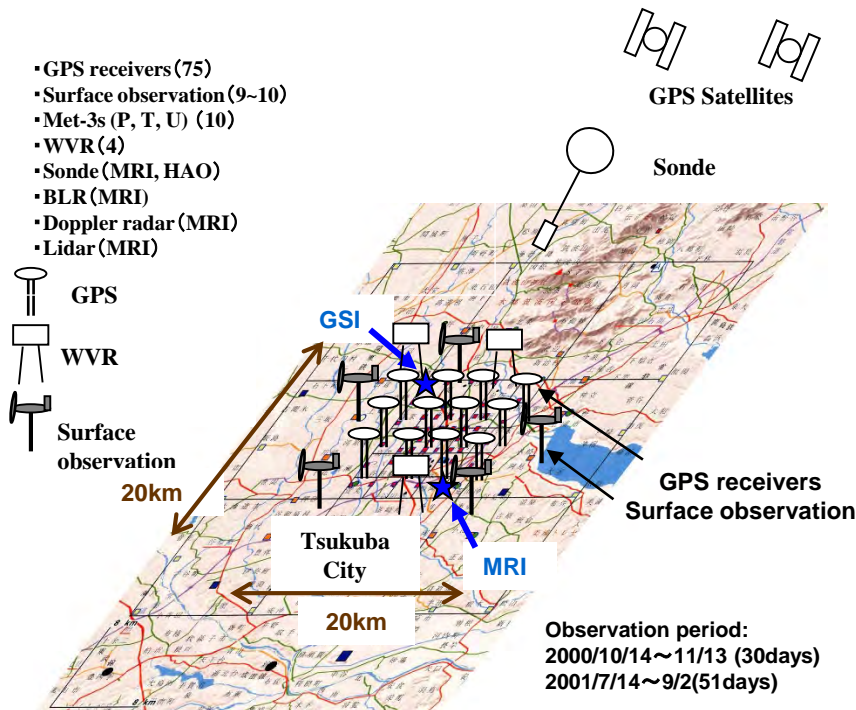
1)Hiromu SEKO* 1)Yoshinori Shoji ,

2)Hajime NAKAMURA and 1)Tetsuya IWABUCHI

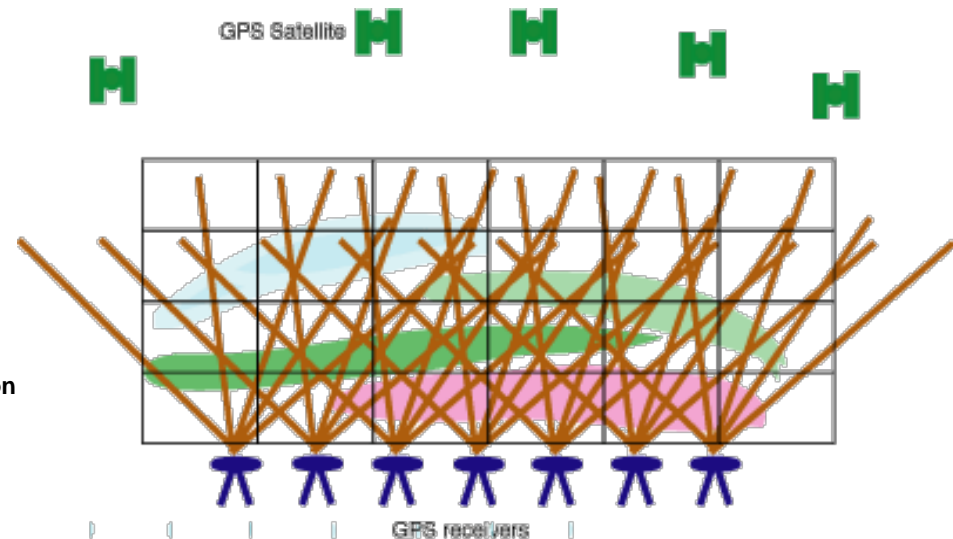
1)Meteorological Research Institute/Japan Meteorological Agency,

2)Numerical Prediction Division/Japan Meteorological Agency

3)Japan Society for the Promotion of Science/Meteorological Research Institute



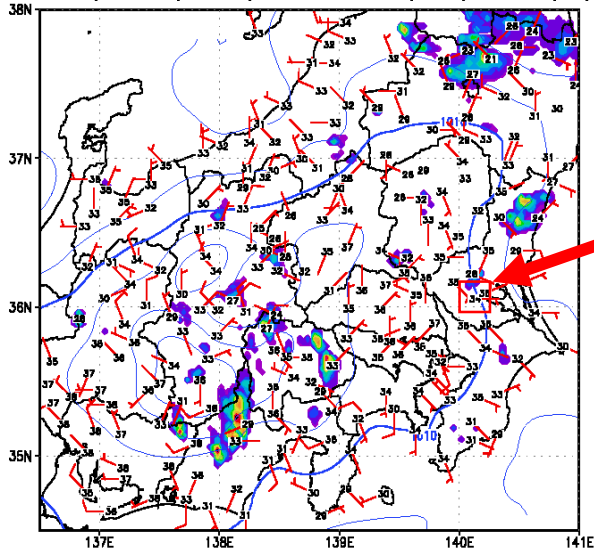
Tsukuba GPS Dense Net



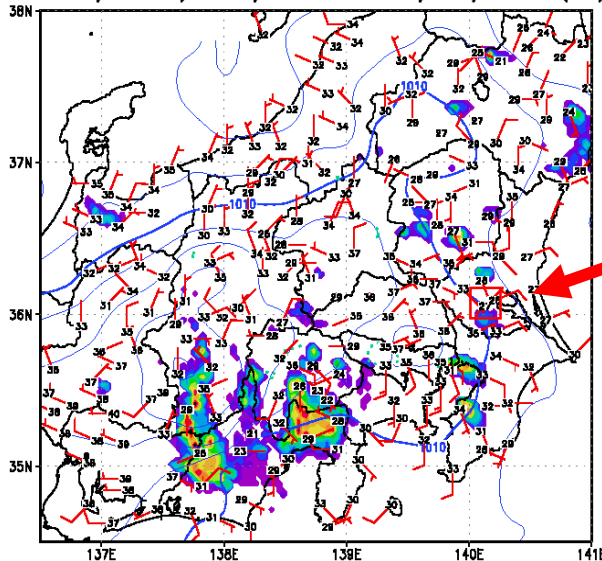
Tomography method

Thunderstorm event on 1st August 2001

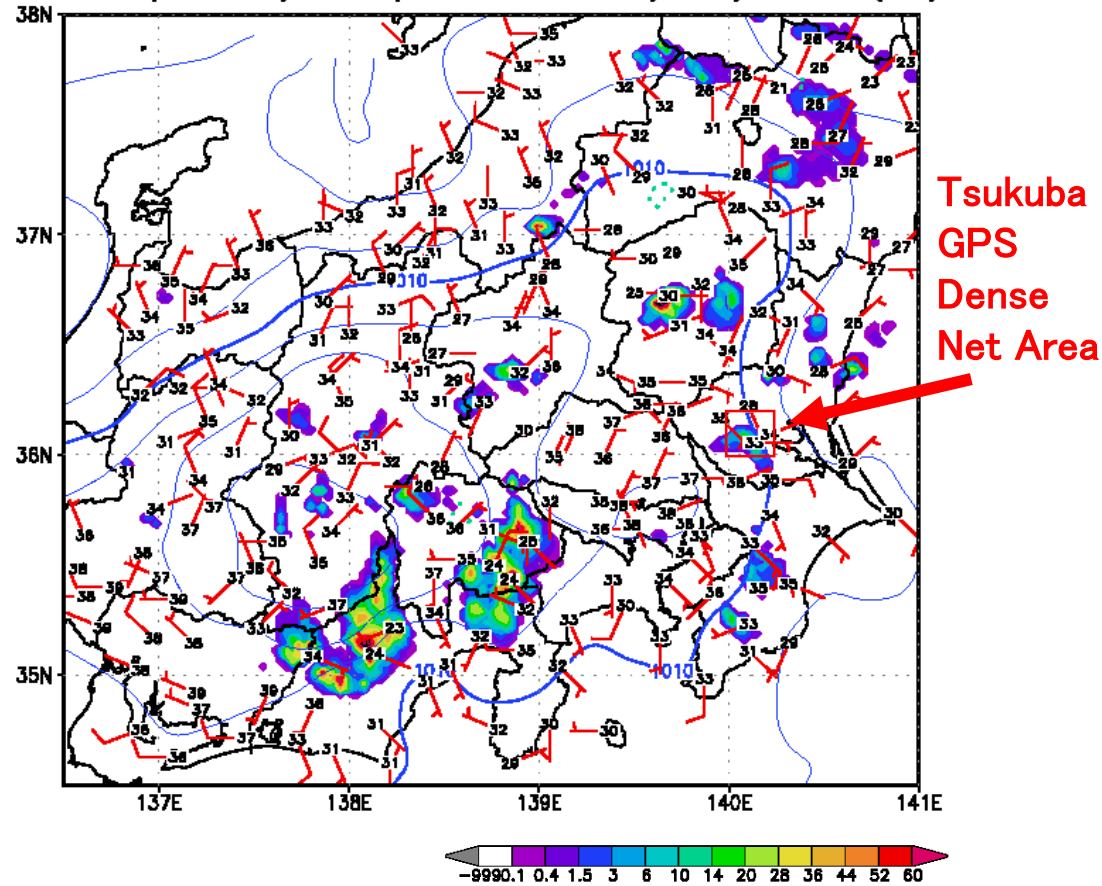
ECHO/Psea/Tsfc/Wsfc 2001/08/01 4(ut)



ECHO/Psea/Tsfc/Wsfc 2001/08/01 6(ut)

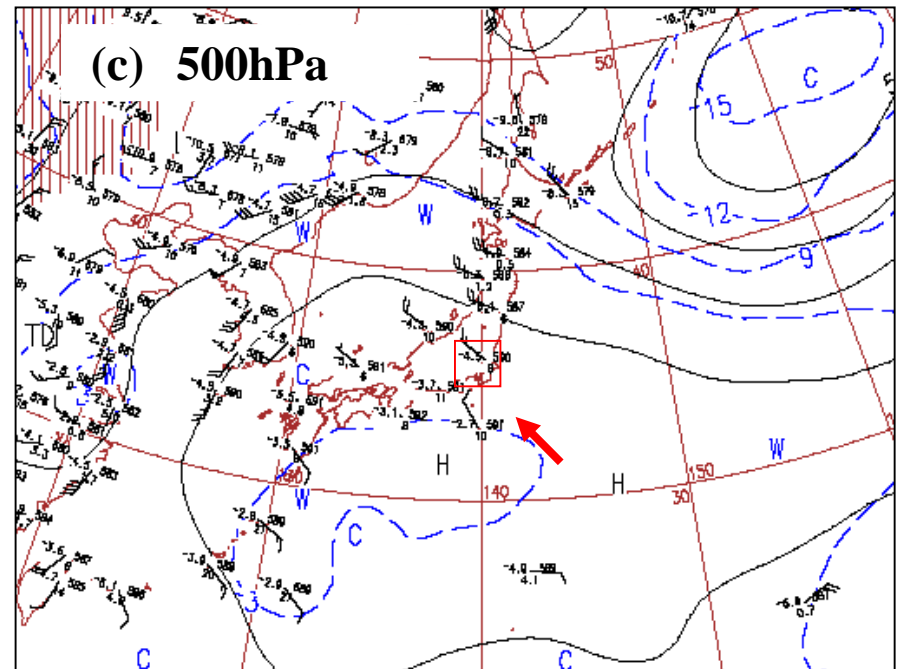
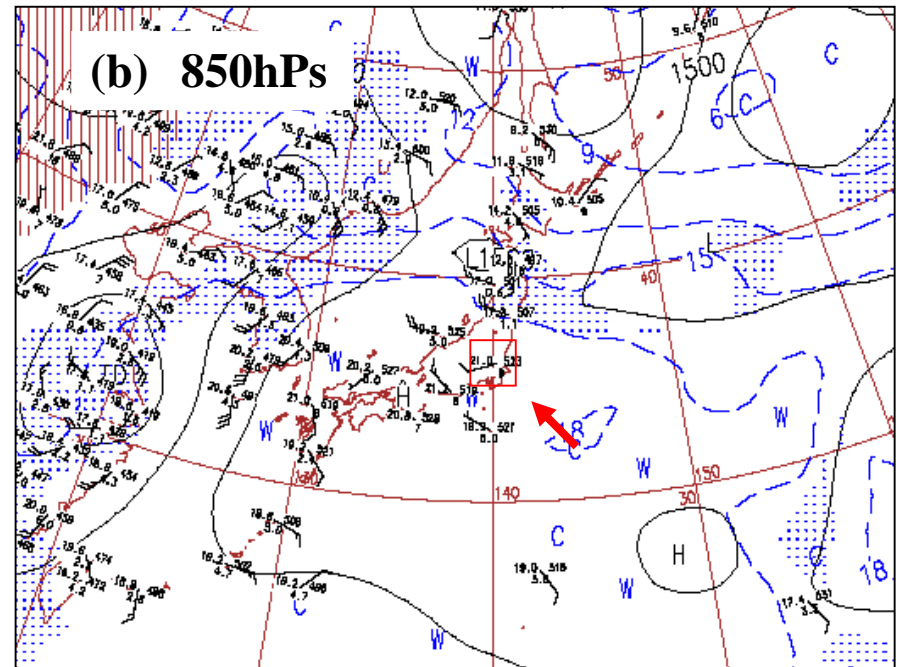
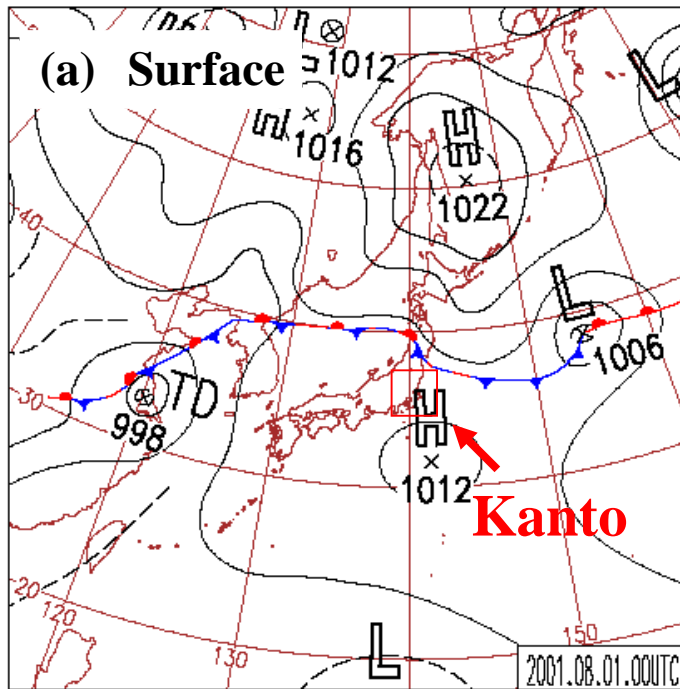


ECHO/Psea/Tsfc/Wsfc 2001/08/01 5(ut)



Tsukuba
GPS
Dense
Net Area

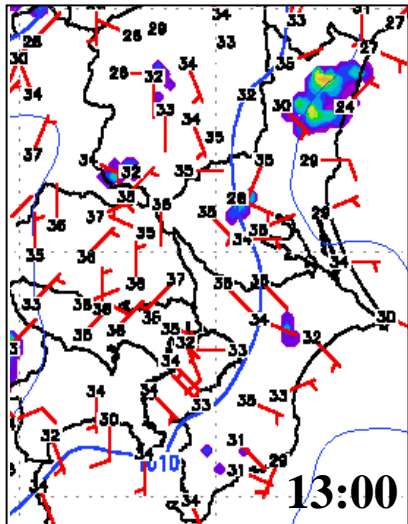
- Small thunderstorm developed at 14JST (=5UTC) and began decaying at 15JST over **Tsukuba GPS Dense net area**.
- Data of GPS Dense Net contains the information of the evolution of this thunderstorm.



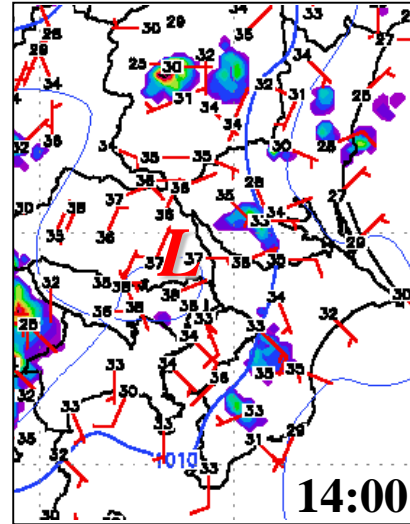
Synoptic scale features

- High pressure area covered **Kanto region**, Japan.
- Thunderstorm was not associated with low pressure or front system.

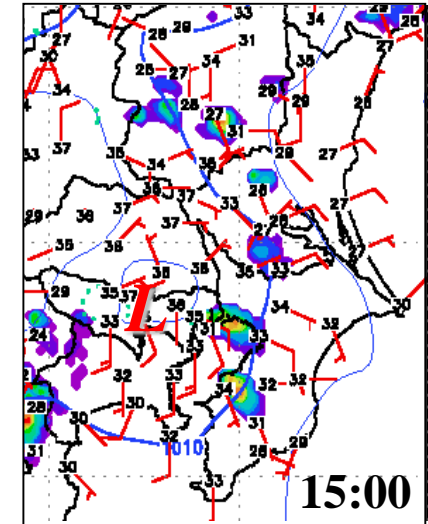
ECHO/Psea/Tsfc/Wsfc 2001/08/01 4(ut)



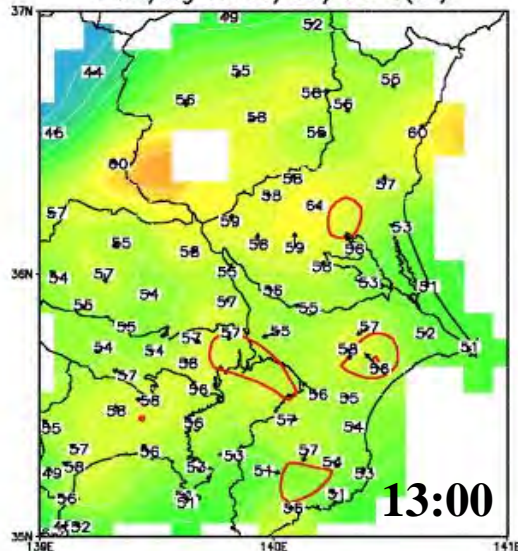
ECHO/Psea/Tsfc/Wsfc 2001/08/01 5(ut)



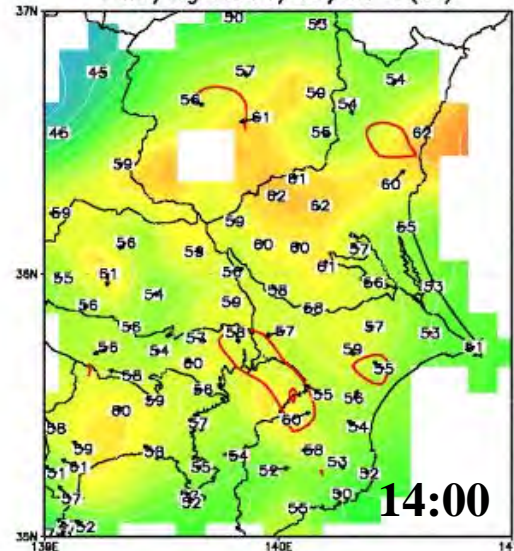
ECHO/Psea/Tsfc/Wsfc 2001/08/01 6(ut)



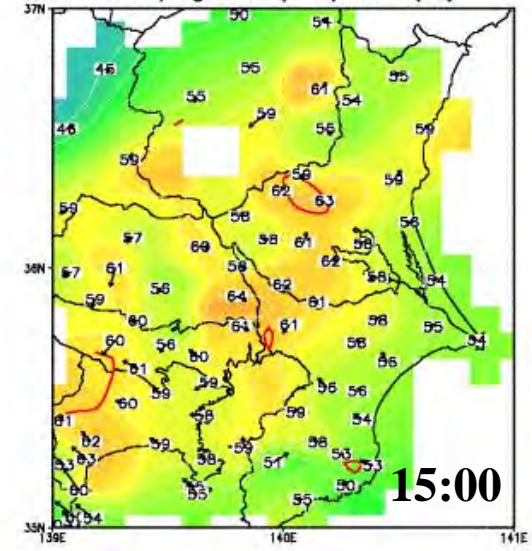
PWV/Lg 2001/08/01 4(ut)



PWV/Lg 2001/08/01 5(ut)

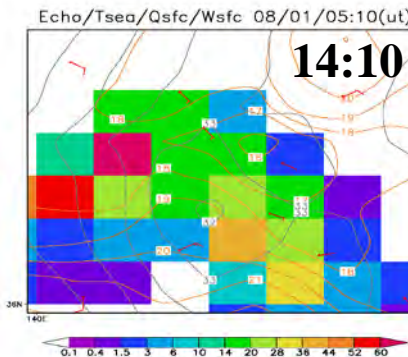
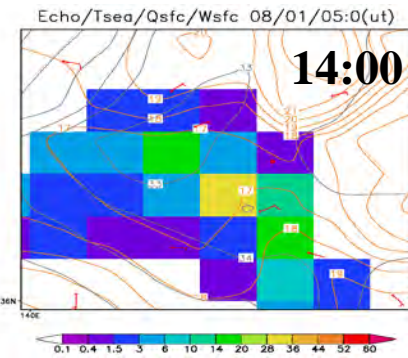
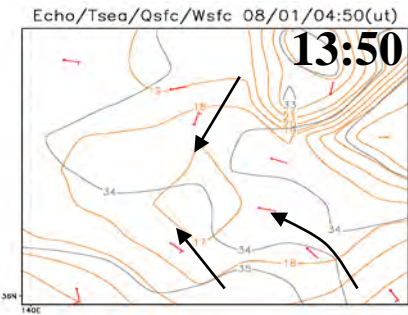


PWV/Lg 2001/08/01 6(ut)

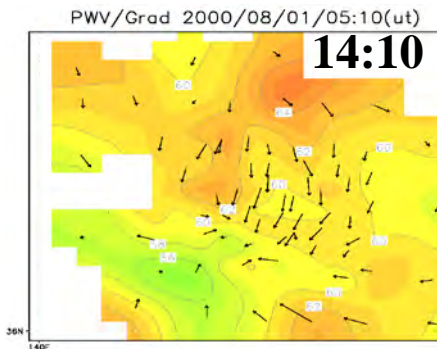
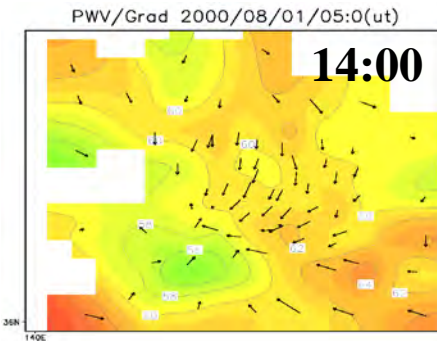
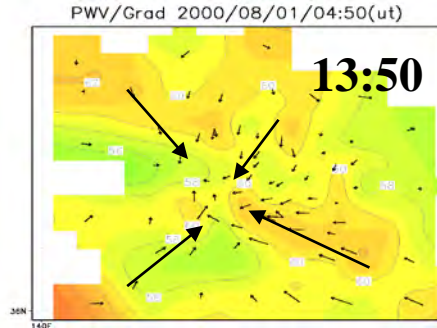


- Large PWV was distributed over the Kanto region
- Water vapor moved toward **Low pressure area** near Tokyo.
- Around Low pressure, large PWV and precipitation areas were generated.

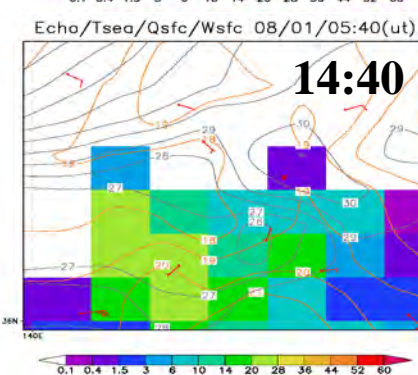
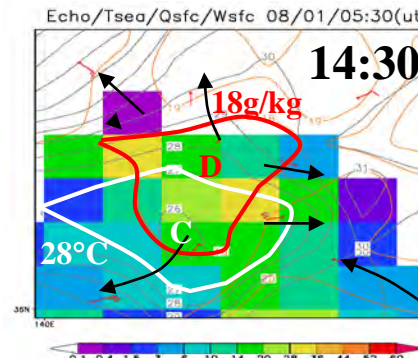
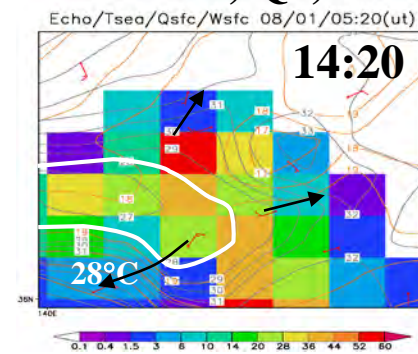
Surface T, Qv, UV



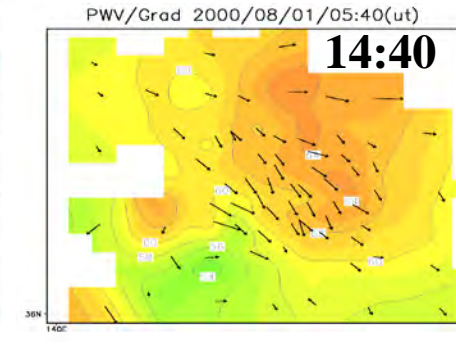
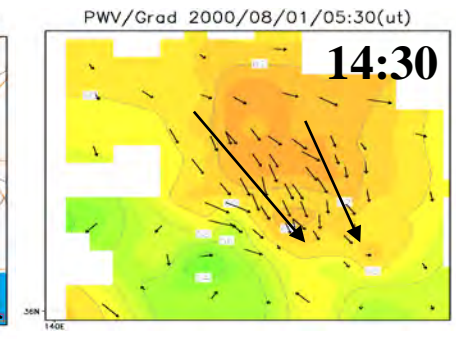
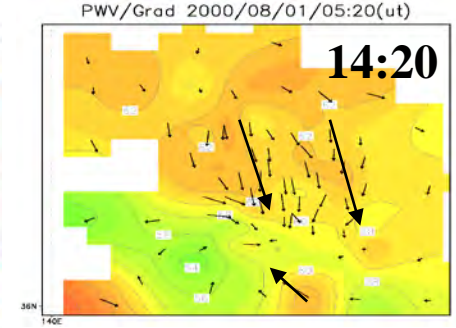
PWV Gradient



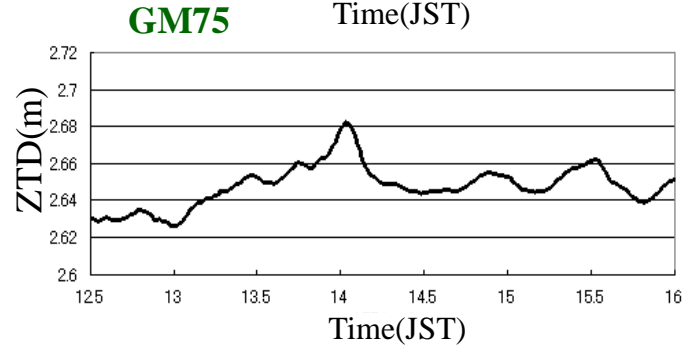
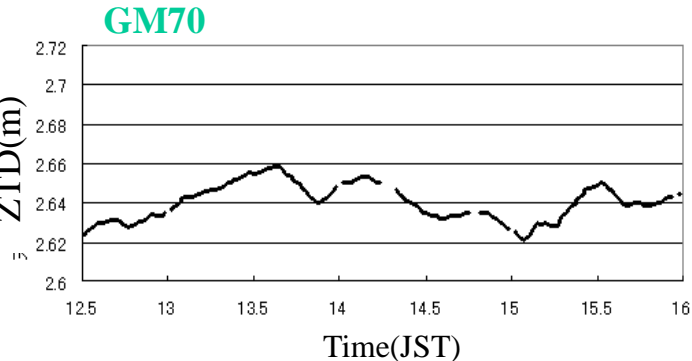
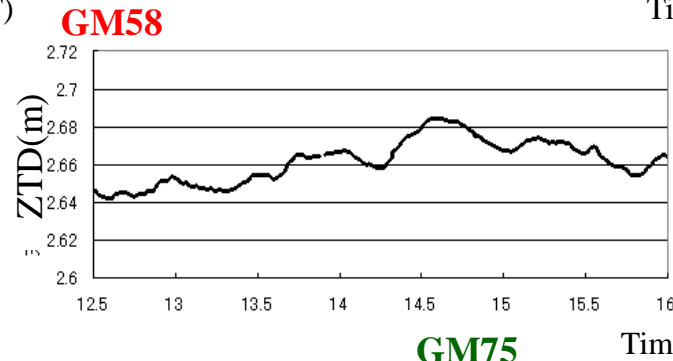
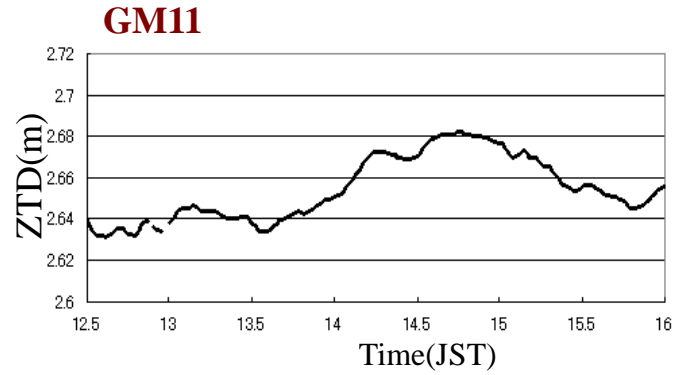
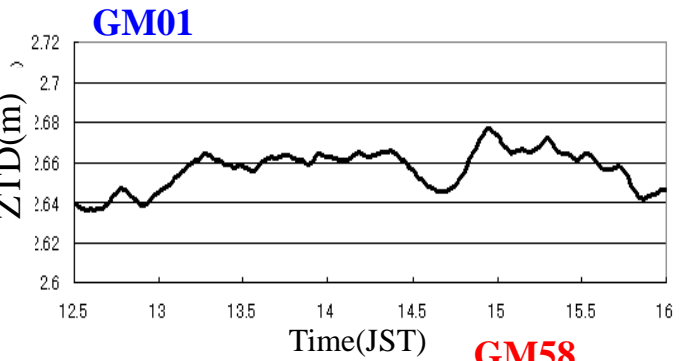
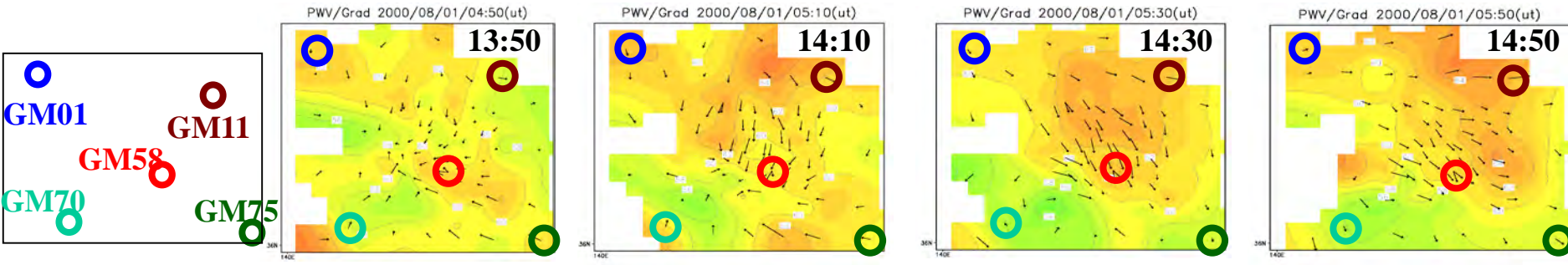
Surface T, Qv, UV



PWV Gradient



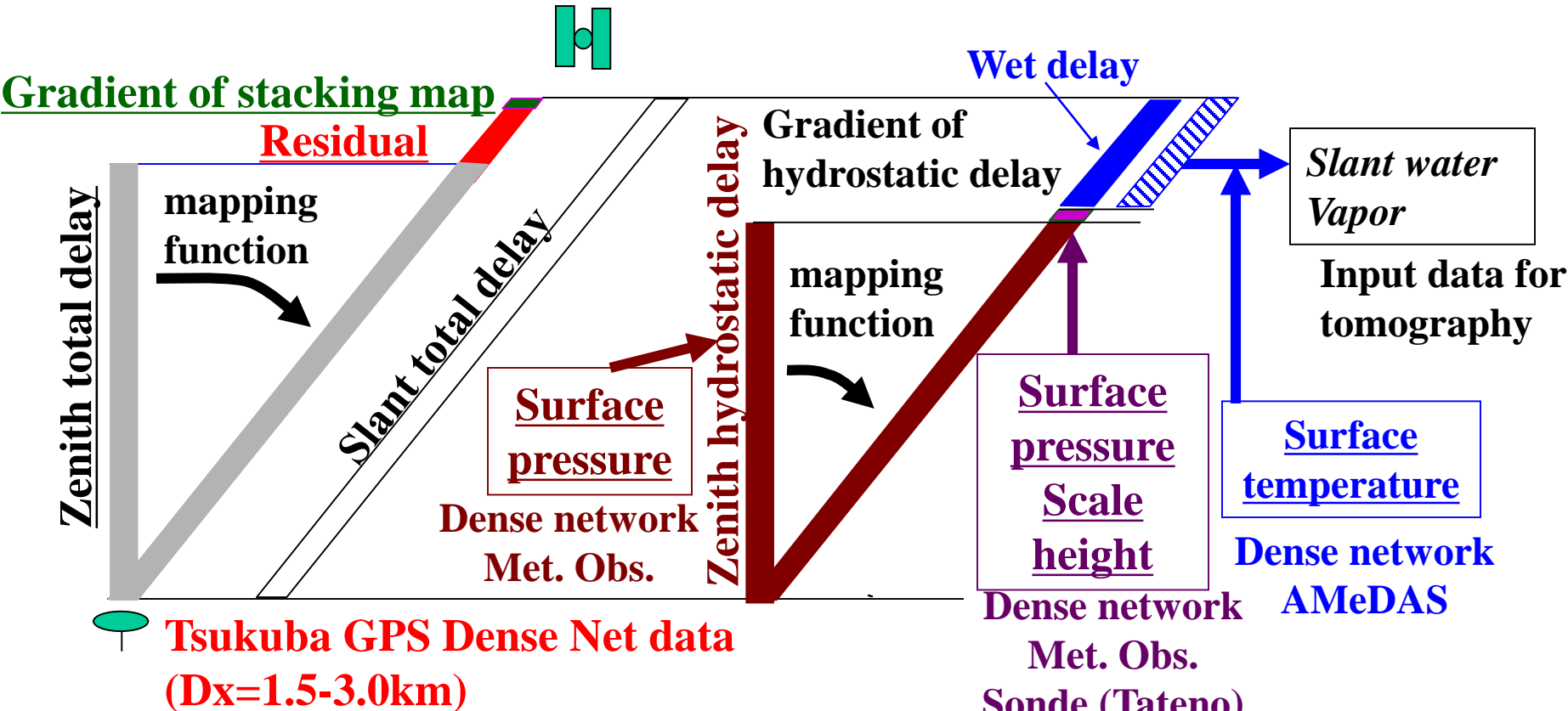
- Precipitation in the southern part decreased surface T, Qv and PWV.
- Surface flows converged (diverged) at 1350JST (1430JST).
- Gradient parameters in the northern part pointed to the south.



■ Variation of ZTD depended on the location of the GPS receivers
 ↓
 Small scale variation of ZTD caused by Thunderstorm.

■ ZTD increased and decreased in the period of 20–40 minutes.

Estimation of slant water vapor (Tsukuba GPS Dense Net)

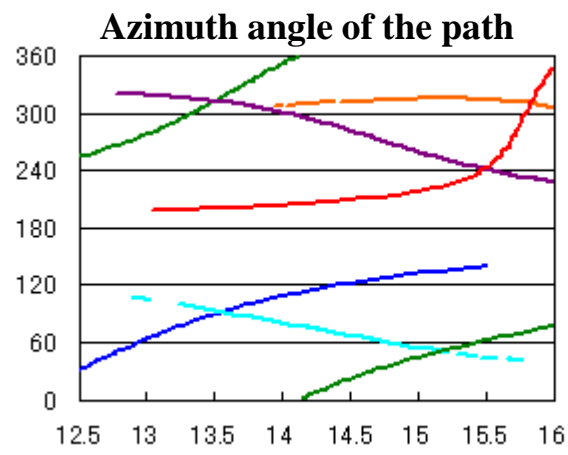
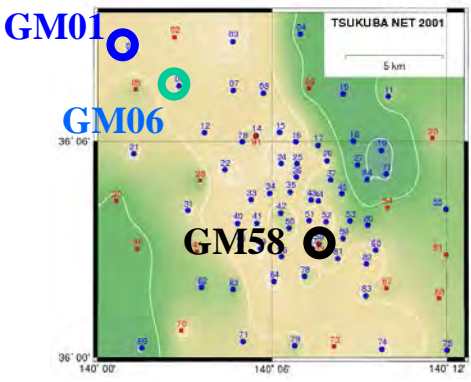


Slant total vapor was estimated from following data:

- Tsukuba GPS Dense Net , Met. Observatory, AMeDAS, Sonde
Automated Meteorological Data Acquisition System

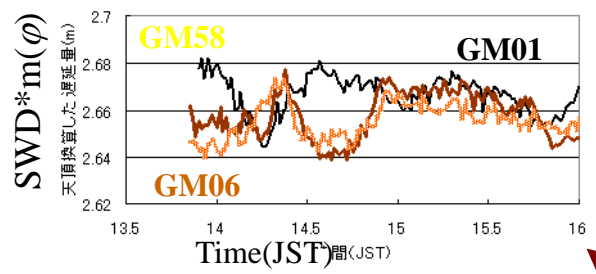
Following effects are taken into consider.

- Gradient of stacking map → climatologic distribution of delay.
- Gradient of hydrostatic delay ← $\nabla_{HP_{surf}} * \text{scale height}$.

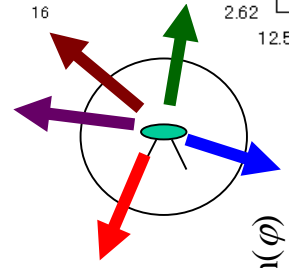
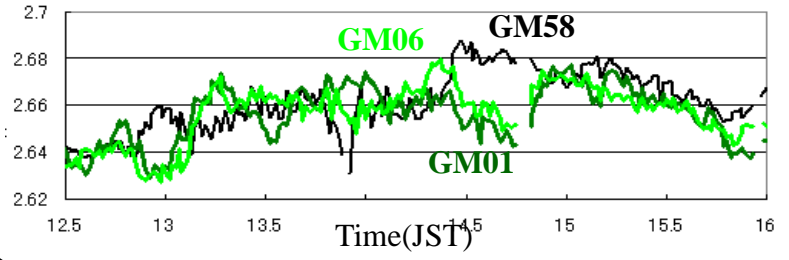


▪ Slant water vapor between the adjacent GPS receivers are correlated.
 → SWV contains the signal of water vapor.

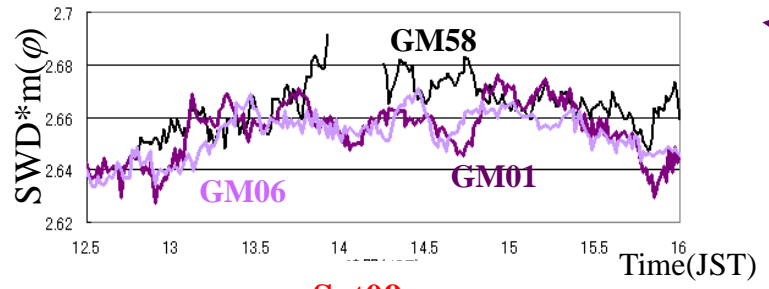
Sat18



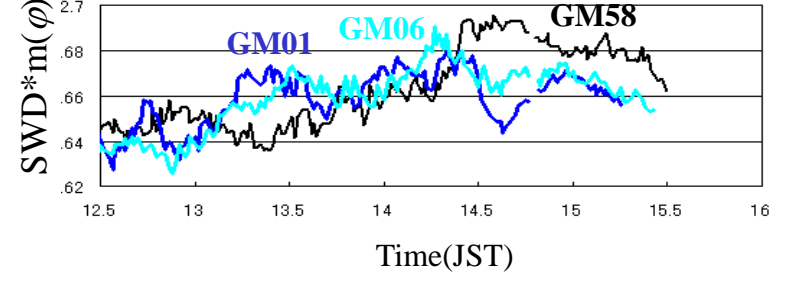
Sat26



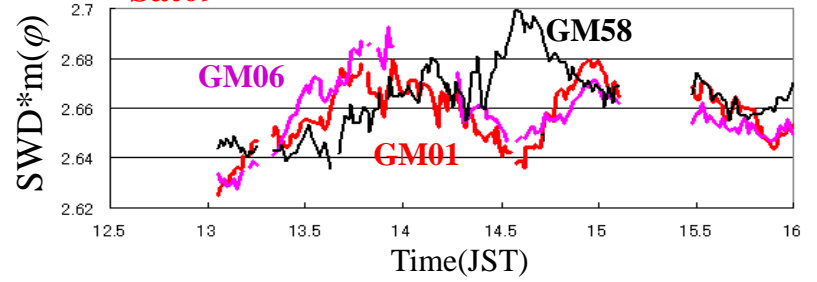
Sat17

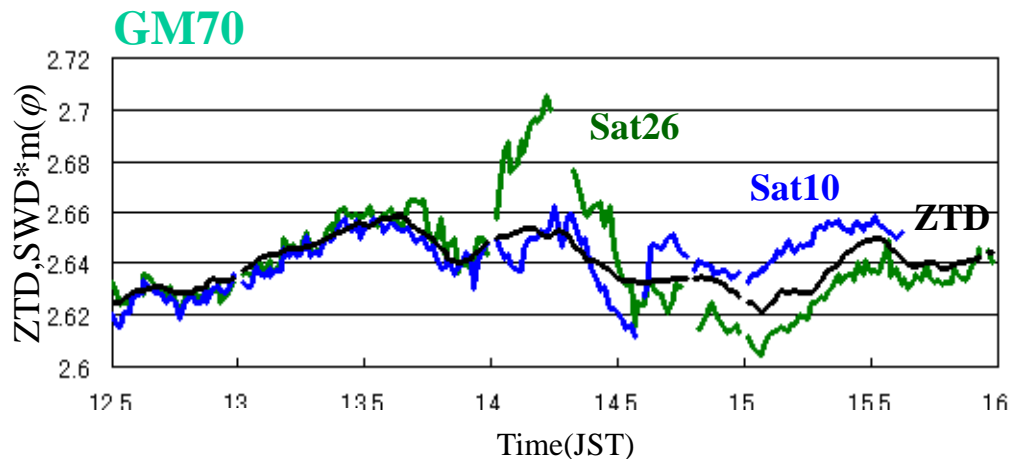
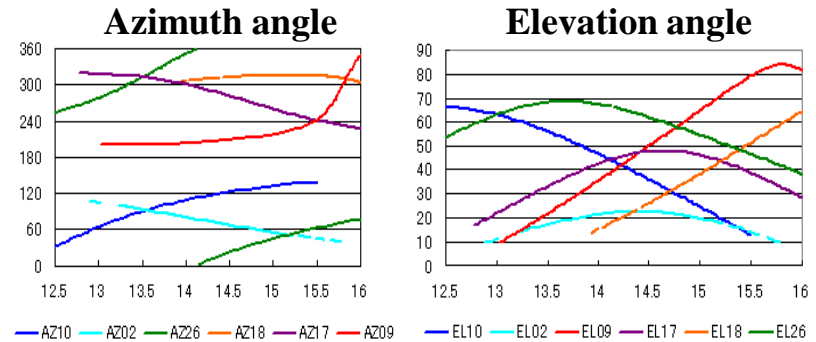
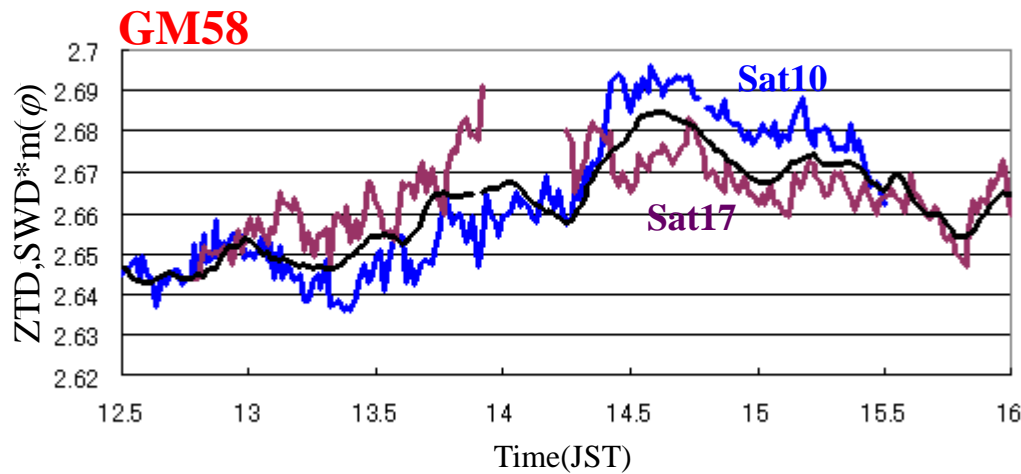
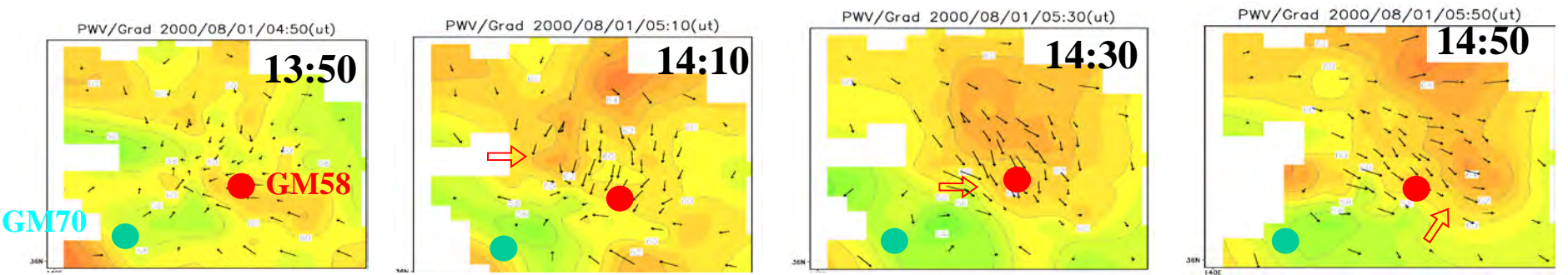


Sat10



Sat09





- Large PWV region (\Rightarrow) passed GM58 at 1430JST.
- SWV of satellite 10 (17) at GM58, which pointed to SE (NW), increased (decreased) at 1430 JST.
- SWV of satellite 26 at GM70, which pointed to north, decreased rapidly at 1430JST.

Method of tomography

Observation equation

- Water vapor along the slant path \mathbf{b}^p is used as input data of tomography method
- Atmosphere is divided into the boxes.

Water vapor is assumed to be uniform in each box (\rightarrow **Box method**)

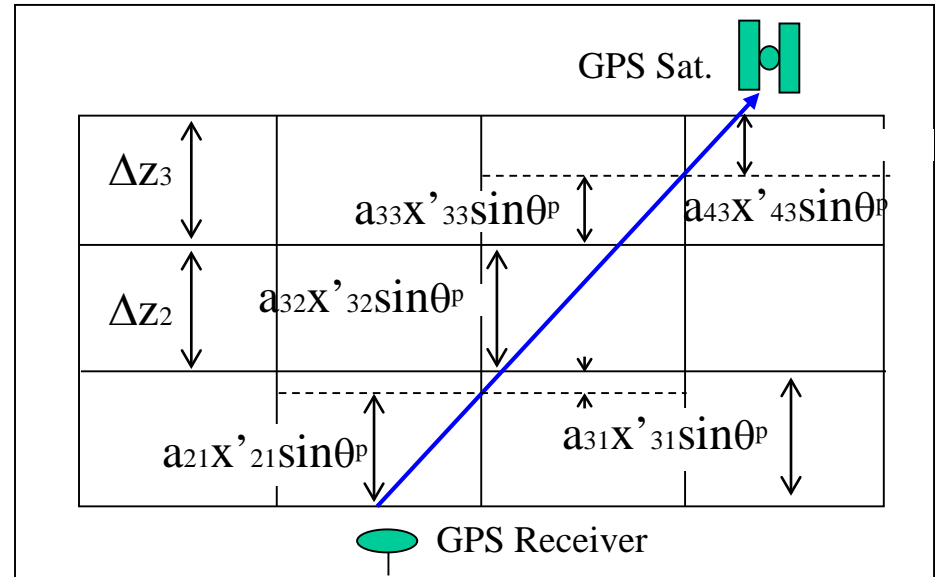
- Water vapor along the slant path \mathbf{b}^p is expressed as follows;

$$\mathbf{a}_{21}\mathbf{X}_{21} + \mathbf{a}_{31}\mathbf{X}_{31} + \mathbf{a}_{32}\mathbf{X}_{32} + \mathbf{a}_{33}\mathbf{X}_{33} + \mathbf{a}_{43}\mathbf{X}_{43} = \mathbf{b}^p,$$

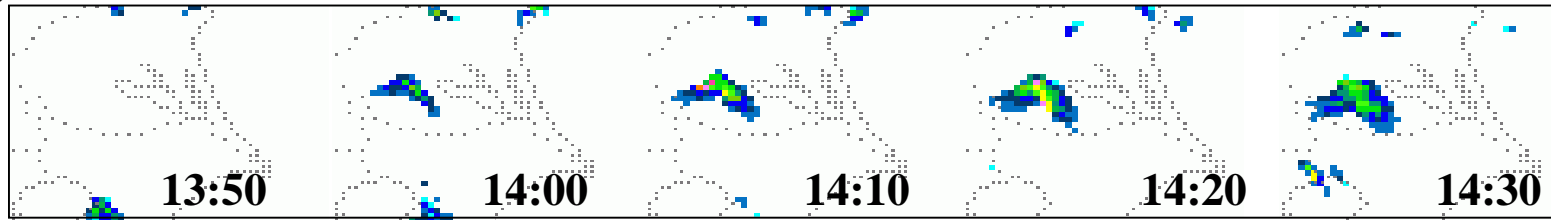
where \mathbf{X}_{ij} is water vapor density and \mathbf{a}_{ij} is path length.

- Path length increases as elevation angle θ^p decreased.

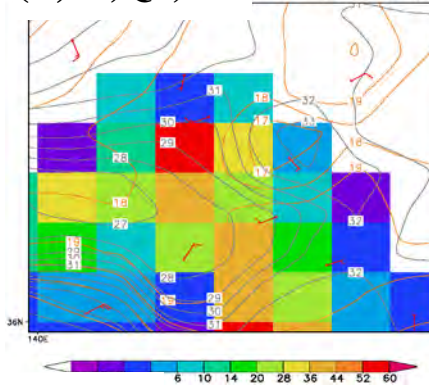
To remove this effects, the equation is multiplied by $\sin\theta^p$



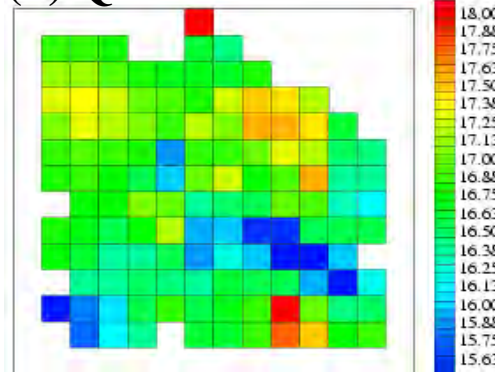
(a) Radar



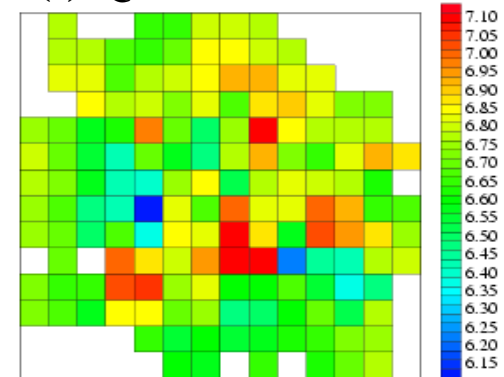
(b) T, Qv, Wind Surface



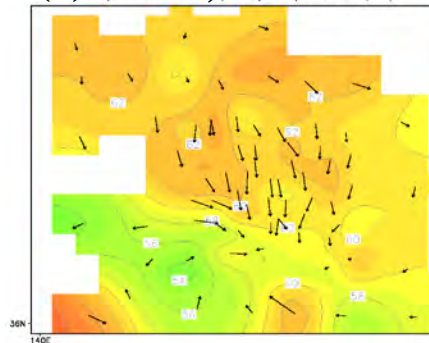
(d) Qv 0-1km



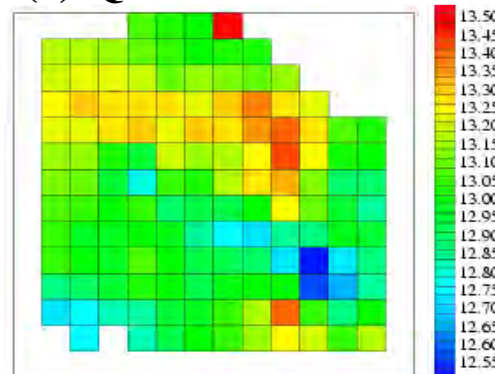
(f) Qv 3-4km



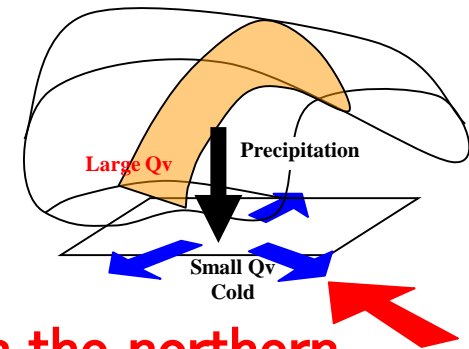
(c) PWV, Gradient



(e) Qv 1-2km



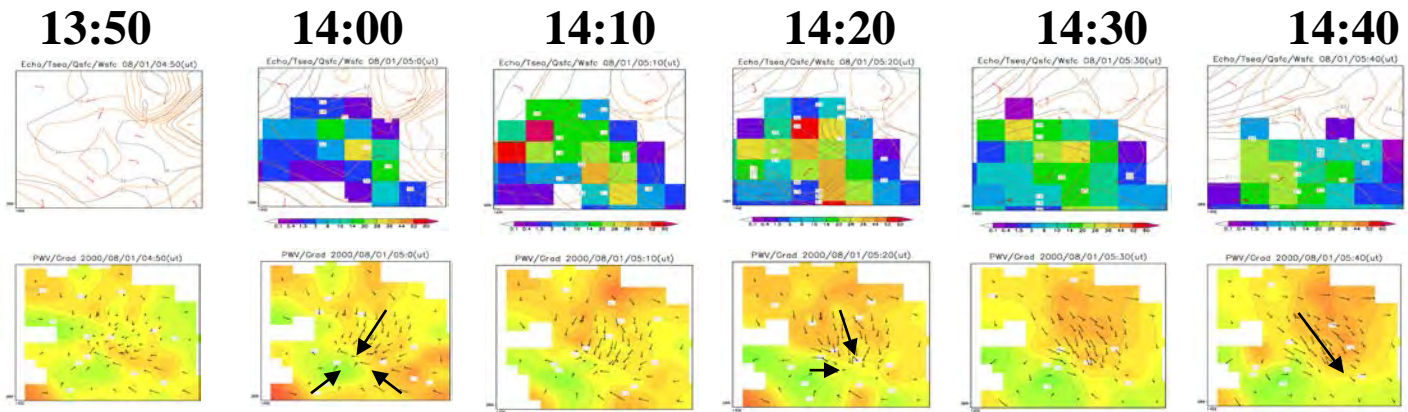
(g) Schematic illustration of precipitation system



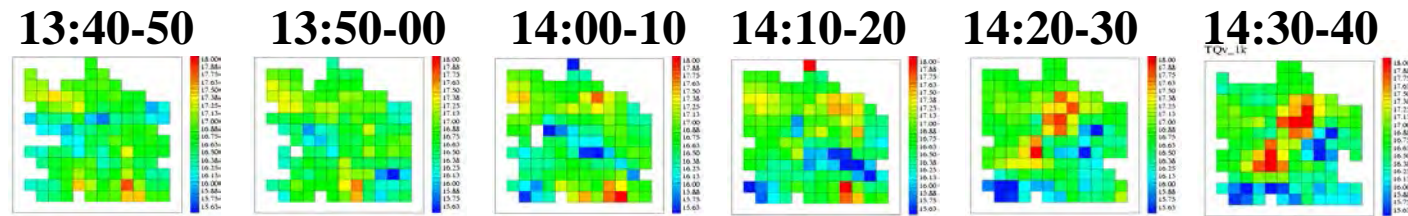
- Region of **large** (small) water vapor existed **on the northern side of precipitation region** (at precipitation region).
- Large water vapor region in the upper layer existed in southern part. This distribution is consistent with gradient parameter.

Surface,
Temp,Qv,
Wind

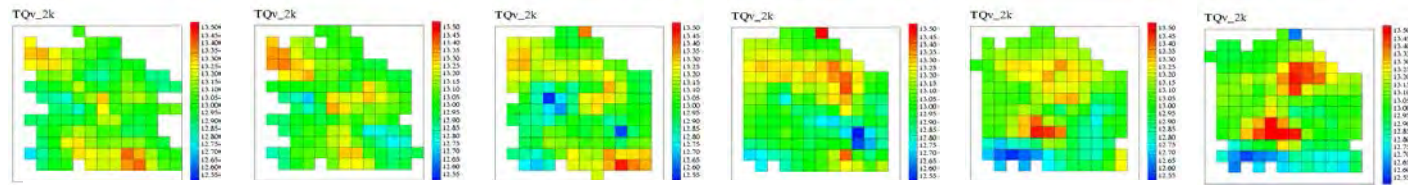
PWV
Gradient



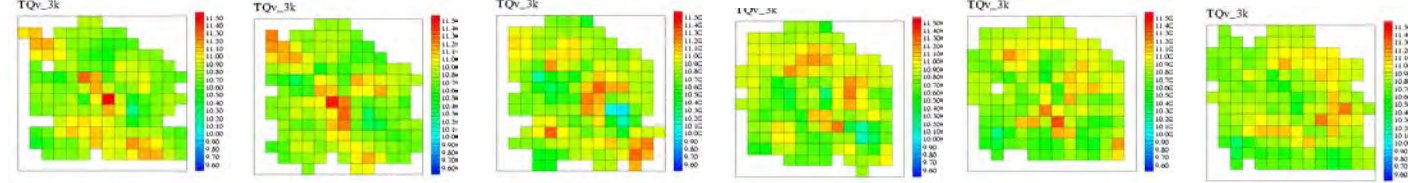
Qv 0-1km



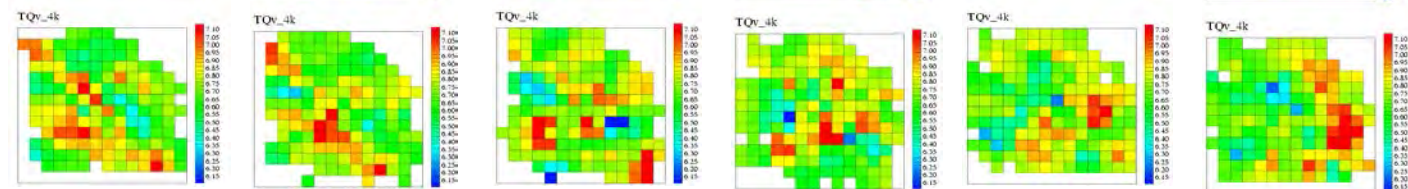
Qv 1-2km



Qv 2-3km



Qv 3-4km



Time variation of water vapor shows that the water vapor deviation did not jump to faraway place.

Summary

- GPS slant water vapor has the signal of the meso- γ scale ($\sim 10\text{km}$) thunderstorm.
- Water vapor variations of the thunderstorm observed by Tsukuba GPS Dense Net:
 - 1) Water vapor decreased under the developed thunderstorm due to the low temperature.
 - 2) Regions of large water vapor at $z > 3.5\text{km}$ correlated to the intense reflectivity regions of 10 minutes later.
 - 3) From 10–20 minutes before the appearance of the echo, water vapor began to increase where the echo would appear, especially $z > 3.5\text{km}$.
 - 4) Most of large water vapor regions at $z > 3.5\text{km}$ correlated to the updraft regions.